



GEOMAGNETICALLY INDUCED CURRENTS IN A BRAZILIAN POWER NETWORK OVER THE SOLAR CYCLES 23 AND 24

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ABSTRACT

Geomagnetically induced currents (GIC) are a space weather effect, which affects ground-based technological structures at all latitudes on the Earth's surface. In this work, we investigate GIC occurrence in a power network at central Brazilian region during the solar cycles 23 and 24. Calculated and measured GIC data are compared for the most intense and moderate geomagnetic storms (*i.e.*, $-150 < \text{Dst} < -50$ nT) of the solar cycle 24. The results obtained from this comparison show a good agreement. The success of the model employed for the calculation of GIC led to the possibility to determine GIC for events during the solar cycle 23 as well. Calculated GIC reached ca. 30 A during the “Halloween storm” in 2003 whilst most frequent intensities lie below 10 A.

Keywords: geomagnetically induced currents, solar cycles, magnetic storm.

Introduction

External geomagnetic field variations induce electric currents in conducting materials at the Earth's surface. These geomagnetically induced currents (GIC) are mostly studied in power network at high-latitude regions due to their frequent hazards. However, recent studies have shown that low to middle-latitude regions may also be subject to problems due to GIC occurrences.

The modeling of GIC in power networks help us to understand the GIC physical processes and can be useful to identify regions more likely to suffer such hazards. GIC calculation requires a model of electrical resistivity in the region of study, direct measurements or model of the geomagnetic field and computation of the geoelectric field and network model of the power grid (Pirjola, 2000; Pirjola, 2002). High-voltage power systems are more vulnerable to GIC flows, especially where they offer low resistance for the electrical current compared to the ground (Radasky, 2011).

In this work, we present a comparison between GIC data measured and calculated in the neutral of a 500 kV transformer at the Itumbiara substation in central Brazil during the solar cycle 24 (Fig. 1). The results from this comparison are used to validate the calculation method, including the power network model. Since the validation succeeds, GIC are calculated for the same power grid considering moderate and intense ($\text{Dst} < -100$) geomagnetic storms during the solar cycle 23.

Calculation of GIC

We applied two steps to calculate GIC at the Itumbiara substation, following the methodology developed by Pirjola (1982): (1) calculation of the geoelectric field and, (2) calculation of the GIC in the power network produced by this geoelectric field following Lehtinen and Pirjola (1985) procedures.

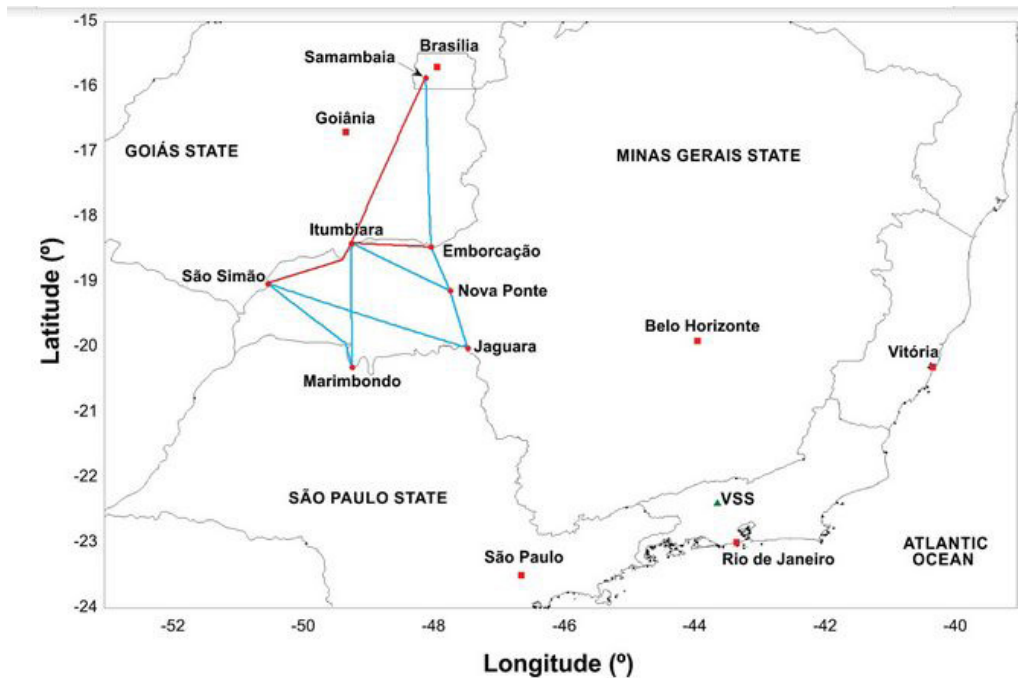
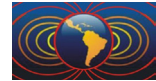


Figure 1. The 500 kV power network investigated in this paper and located in central Brazil. The transmission lines shown in blue were installed in 2006.

Calculation of the geoelectric field

We used the plane wave model in Cartesian coordinate system (Cagniard, 1953; Pirjola, 1982) to calculate the geoelectric field (GEOF). In the plane wave approximation, the time series corresponding to geomagnetic data is Fourier transformed to the frequency domain, and then the geoelectric northward (E_x) and eastward (E_y) components are obtained from the following relation:

$$E_{x,y} = \pm \frac{Z(\omega)B_{y,x}}{\mu_0} \quad (1)$$

Where $Z(\omega)$ is the surface impedance in function of the (angular) frequency ω and on the Earth's conductivity structure, and $B_{y,x}$ are the geomagnetic eastward and northward components, respectively. Once the E_x and E_y components are obtained in the frequency domain from Eq. (1), they are converted to the time domain by the application of an inverse Fourier transform.

In this study, we use a ground conductivity model obtained from magnetotelluric surveys, as described by Bologna et al. (2001). It consists of four layers, with the deepest layer being infinitely thick, and approximates to the local geology in the area of the 500 kV power network shown in Figure 1.

The magnetic observatory closest to the Itumbiara substation is located in Vassouras, Rio de Janeiro State, (VSS, 22.4°S, 43.6°W, and see Figure 1). The distance between VSS and Itumbiara is about 740 km considered reasonable to use VSS dataset in this work since geomagnetic field morphology and time variations are not so different between VSS and Itumbiara substation (Trivedi *et al.*, 2007).

GIC modelling using the Lehtinen-Pirjola (LP) method

The (geo)voltage between two points i and j at the Earth's surface, can be calculated by integrating GEOF along the trajectory, which, is the transmission line between i and j :



$$V_{ij} = \int_i^j \mathbf{E} \cdot d\mathbf{s} \quad (2)$$

The power grid is modelled as a network of N earthed substations (*i.e.*, nodes) interconnected by transmission lines with voltage sources in series with the line resistances. In the LP method, the GIC flowing into the ground at the nodes are expressed as an $N \times 1$ column matrix I_e (Lehtinen and Pirjola 1985):

$$I_e = (Id + Y_n Z_e)^{-1} J_e \quad (3)$$

where Id is the $N \times N$ identity matrix, Y_n is the $N \times N$ network admittance matrix with their elements given by:

$$Y_{ij} = \frac{-1}{R_{ij}} \quad (i \neq j); \quad Y_{ii} = \sum_{i \neq k}^N \frac{1}{R_{ik}} \quad i = j \quad (4)$$

R_{ij} is the resistance of the line between nodes i and j and Z_e is the $N \times N$ earthing impedance matrix with the diagonal elements being the earthing resistances $(r_i)(r_i)$ of the nodes:

$$Z_{ii} = r_i \quad (5)$$

If the earthing points are far apart (as in case of different substations), the off-diagonal elements of Z_e are zero. The elements of the $N \times 1$ column matrix J_e are defined by:

$$J_e = \sum_{i=1; i \neq e}^N \frac{V_{ie}}{R_{ie}} \quad (j = 1, \dots, N) \quad (6)$$

J_e thus includes the information about the geovoltages V_{ij} . Since $I_e = J_e$ in the case of perfect earthings of the nodes, (*i.e.*, $Z_e = 0$), the elements of J_e can be interpreted as “perfect earthing currents”.

Results and discussion

In order to compare the measured and calculated GIC results, we choose three most intense geomagnetic storm events. They are included in the panels (a) 16 June 2012, (b) 2 October 2013 and (c) 8 October 2013 of Figure 2, where the blue and red curves depict the calculated and measured GIC, respectively. The peaks have a good agreement in the calculated and measured GIC data. Regarding specifically the event on 2 October 2013 (Figure 2b), the curves are also similar when the GIC values are smaller, so a good agreement between calculations and measurements holds true most of the time. We suggest that the small discrepancies at low amplitudes of GIC can be related to local ionospheric instabilities. The general behavior observed in comparison between calculated and measured GIC data indicates that the calculations are reliable and can thus be applied to the determination of GIC at the Itumbiara substation during periods with no measurements available.

The GIC for large magnetic storms ($Dst < -100$ nT, total of 90 events) that occurred during solar cycle 23 (listed in Pandey, 2009) was calculated for the day when the Dst reaches its minimum value, the day previous and the day after. Figure 3 shows the daily maximum amplitudes of calculated GIC versus maximum decreases in Dst value. We observe that several storms (moderate intensity) can cause GIC values higher than 10 A and a few (intense) storms can also lead to more than 20 A. The highest GIC amplitudes obtained at the Itumbiara power network, during solar cycle 23 was 30 A, during the Halloween storm ($Dst = -401$ nT). However, to illustrate the complexity of the phenomena involved in the GIC generation, it is observed in Figure 3 two geomagnetic storms producing Dst higher than -400 nT. For these two storms were observed GIC values ca. 30 A and 15 A, respectively. This important difference is observed because GIC could have other contributions that are not properly described by Dst index amplitude.

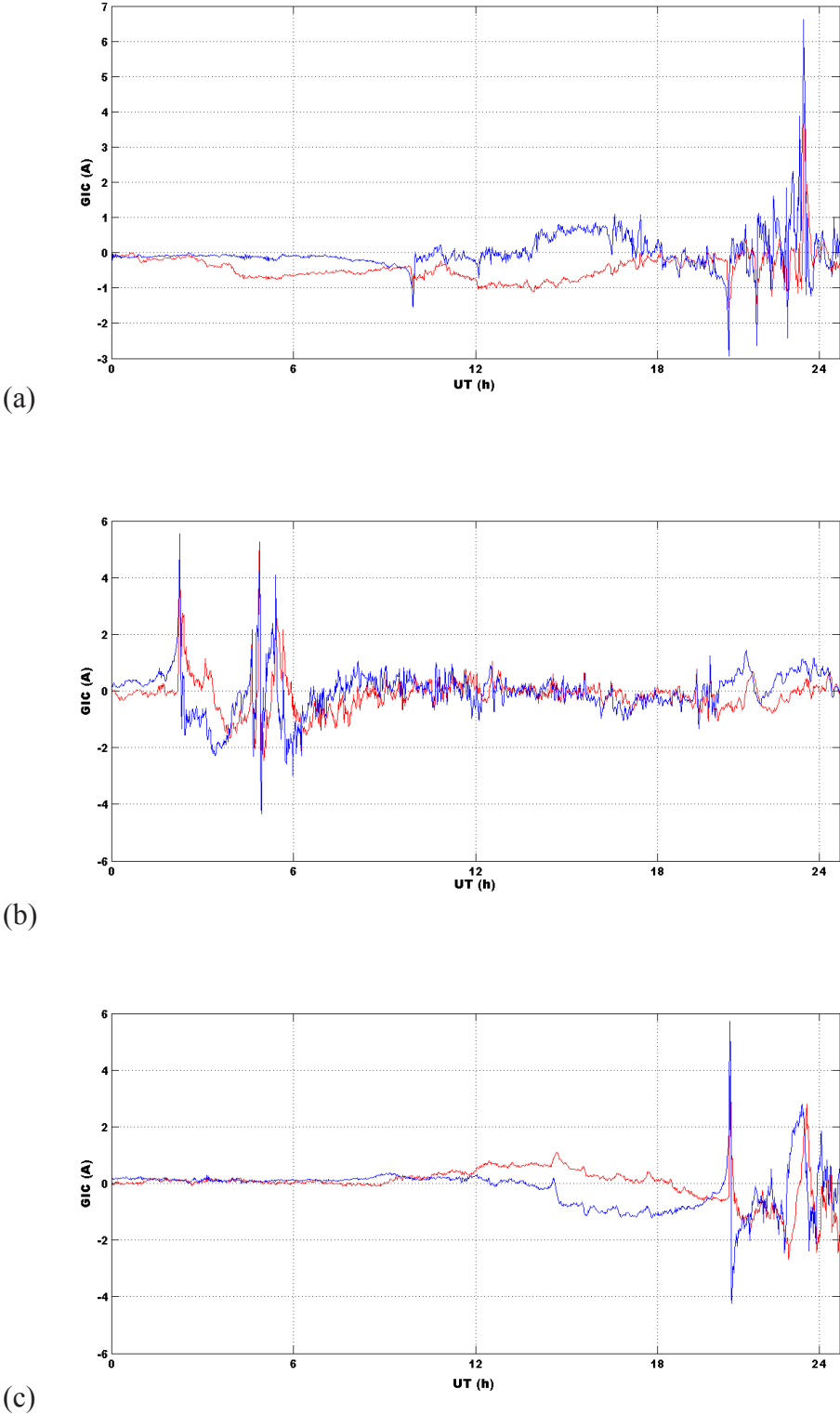


Figure 2. Calculated (blue) and measured (red) GIC in the neutral of the Itumbiara 500 kV transformer on (a) 16 June 2012, (b) 2 October 2013, (c) 8 October 2013.

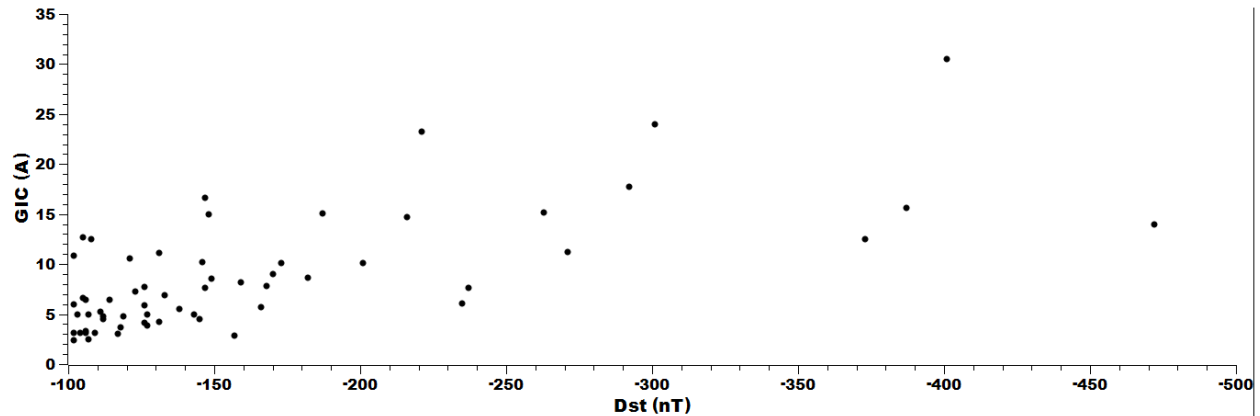


Figure 3. Daily maximum amplitude of GIC in the neutral of the Itumbiara 500 kV transformer versus the intensity of the magnetic storm on solar cycle 23.

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